



Balance of
wind-driven
seas

Kinetic
equation

The
misguiding
star

Nonlinear
forcing and
damping

An
example:
Mixed sea

Summary

Scales of nonlinear relaxation and the problem of balance of wind-driven seas

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Klauss HASSELMANN 1962, On the nonlinear energy transfer in a gravity wave spectrum

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Kinetic equation

$$\frac{\partial n_{\mathbf{k}}}{\partial t} + \nabla_{\mathbf{k}} \omega_{\mathbf{k}} \nabla_{\mathbf{r}} n_{\mathbf{k}} = \mathcal{S}_{in} [n_{\mathbf{k}}] + \mathcal{S}_{diss} [n_{\mathbf{k}}] + \mathcal{S}_{nl} [n_{\mathbf{k}}]$$

$n(\mathbf{k})$ – spatial spectrum of wave action (Fourier amplitudes squared for deep water $kh \gg 1$)

Important! Right-hand side

- Wave input \mathcal{S}_{in} – empirico-heuristical
- Dissipation \mathcal{S}_{diss} – empirico-heuristical
- Nonlinear transfer \mathcal{S}_{nl} – from “the first principles”



Komen, Hasselmann, Hasselmann 1984, "On the existence of a fully developed wind-sea spectrum"

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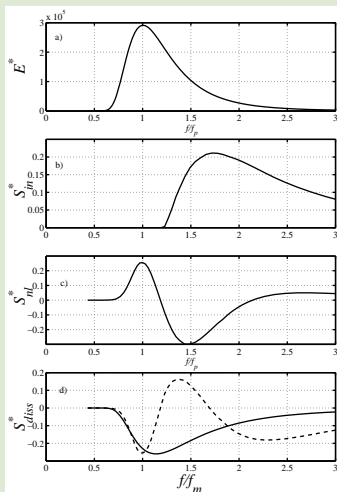
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- Pierson-Moskowitz spectrum
- Input by Snyder et al. (1981)

$$S_{in}(\omega) = \max(0, 0.25 \rho_a / \rho_w \omega^3 \times (28 u_* / C_p - 1));$$

- Dissipation

$$S_{diss} = -S_{in} - S_{nl}$$

Output

$$\tilde{S}_{in} : \tilde{S}_{nl} : \tilde{S}_{diss} = 3 : (-1) : (-2)$$

$$\tilde{S}_i = \int_0^{2.5 f_m} S_i df d\theta$$



Q. What is responsible for wind-wave balance?

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Answers

- 1 Mainstream
Wave input and dissipation provide a relaxation to an inherent state
- 2 Non-conventional ?
Conservative nonlinear transfer term contains both forcing and damping and is able to provide the strong relaxation on its own !!!



Nonlinear forcing and damping

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$$S_{nl} = \pi g^2 \int |T_{0123}|^2 (N_1 N_2 N_3 + N N_{\mathbf{k}_2} N_{\mathbf{k}_3} - N N_{\mathbf{k}_1} N_{\mathbf{k}_2} - N N_{\mathbf{k}_1} N_{\mathbf{k}_3}) \times \delta(\mathbf{k} + \mathbf{k}_1 - \mathbf{k}_2 - \mathbf{k}_3) \delta(\omega_{\mathbf{k}} + \omega_1 - \omega_2 - \omega_3) d\mathbf{k}_1 d\mathbf{k}_2 d\mathbf{k}_3 \quad (1)$$

Split into two terms

$$S_{nl} = F_{\mathbf{k}} - \Gamma_{\mathbf{k}} N_{\mathbf{k}} \quad (2)$$

where

$$F_{\mathbf{k}} = \pi g^2 \int |T_{0123}|^2 N_1 N_2 N_3 \times \delta(\mathbf{k} + \mathbf{k}_1 - \mathbf{k}_2 - \mathbf{k}_3) \delta(\omega_{\mathbf{k}} + \omega_1 - \omega_2 - \omega_3) d\mathbf{k}_1 d\mathbf{k}_2 d\mathbf{k}_3 \quad (3)$$

$$\Gamma_{\mathbf{k}} = \pi g^2 \int |T_{0123}|^2 (N_1 N_2 + N_1 N_3 - N_2 N_3) \times \delta(\mathbf{k} + \mathbf{k}_1 - \mathbf{k}_2 - \mathbf{k}_3) \delta(\omega_{\mathbf{k}} + \omega_1 - \omega_2 - \omega_3) d\mathbf{k}_1 d\mathbf{k}_2 d\mathbf{k}_3 \quad (4)$$



Split S_{nl} into two terms (N.N. Ivenskikh approach based on WRT-algorithm)

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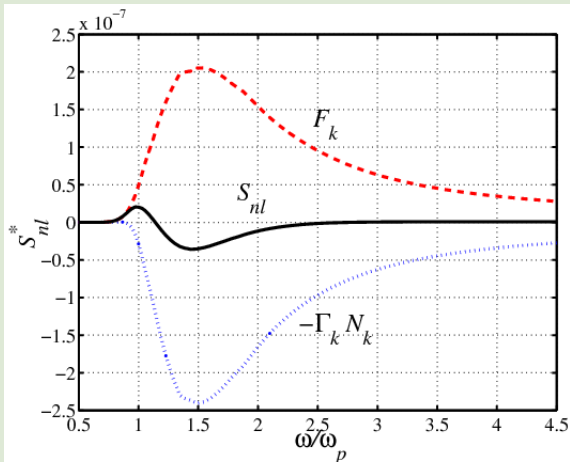
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S_{nl} is small due to proximity to an inherent state !



Theoretical estimate of Γ_{nl}

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Hypothesis: The key contribution to Γ_{nl} is from interactions of pairs of long and short waves

$$\Gamma_{\mathbf{k}} \approx 2\pi g^2 \int |T_{0103}|^2 \delta(\omega_0 - \omega_3) N_1 N_3 d\mathbf{k}_1 d\mathbf{k}_3 \quad (5)$$

$$T_{0103} \approx 2|\mathbf{k}_1|^2 |\mathbf{k}| \cos \Theta$$

$$\Gamma_{\mathbf{k}} = 36\pi\omega (\omega/\omega_p)^3 \mu_p^4 \cos^2 \Theta, \quad (6)$$

small parameter $\mu_p = \sqrt{\frac{E\omega_p^4}{g^2}}$ – wave steepness

An enhancing factor: $36\pi \approx 113.1$



Compare nonlinear damping decrement and wind input increment

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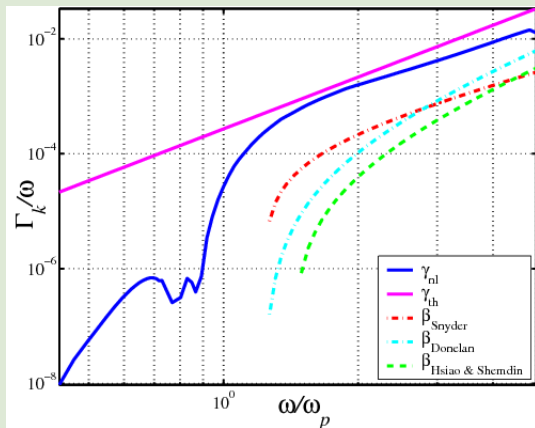
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S_{nl} surpasses S_{in} and S_{diss} in order of magnitude !



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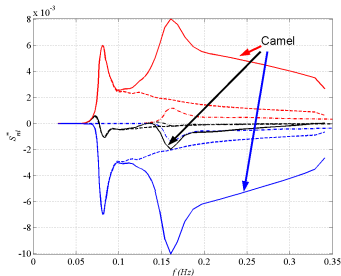
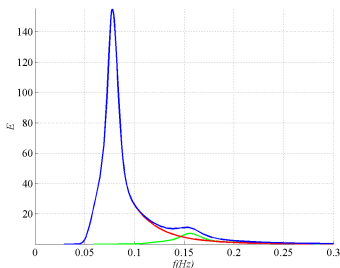
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How to use our simple estimate of nonlinear damping?

$$\Gamma_k = 36\pi\omega(\omega/\omega_p)^3 \mu_p^4 \cos^2 \Theta,$$

High (ω/ω_p) – from swell peak, high μ_p – from wind waves

I. Young, 2006, JGR



See Badulin et al. 2008, Rogue Waves 2008



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- 1 **Nonlinear relaxation**, generally, is much stronger than quasi-linear external forcing and wave dissipation;
- 2 **Interactions of long and short waves** play key role in this relaxation;
- 3 **We do not ignore wave input and dissipation**, we just put them into proper place