



Balance of
wind-driven
seas

Kinetic
equation

Wave input
Nonlinear transfer

The
misguiding
star

Nonlinear
forcing and
damping

An example:
Mixed sea

A model of the
mixed sea

Summary

ONCE MORE ON THE DOMINANCE OF NONLINEAR TRANSFER

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Outline

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- 1 Kinetic (Hasselmann's) equation for wind-driven waves
 - Wave input
 - Nonlinear transfer
- 2 The misguiding star by Komen, Hasselmann, Hasselmann 1984
- 3 Nonlinear forcing and damping
- 4 An example: Mixed sea
 - “The least sophisticated” model of the mixed sea
- 5 Summary

You are welcome to copy this presentation



Klauss HASSELMANN 1962, "On the nonlinear energy transfer in a gravity wave spectrum"

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Kinetic equation

$$\frac{\partial n_{\mathbf{k}}}{\partial t} + \nabla_{\mathbf{k}} \omega_{\mathbf{k}} \nabla_{\mathbf{r}} n_{\mathbf{k}} = S_{in} [n_{\mathbf{k}}] + S_{diss} [n_{\mathbf{k}}] + S_{nl} [n_{\mathbf{k}}]$$

$n(\mathbf{k})$ – spatial spectrum of wave action (Fourier amplitudes squared for deep water $kh \gg 1$)

Important! Right-hand side

- Wave input S_{in} – empirico-heuristical
- Dissipation S_{diss} – empirico-heuristical
- Nonlinear transfer S_{nl} – from “the first principles”



Wave input. The Vavilov-Cherenkov excitation Cherenkov, Frank, Tamm – Nobel Prize 1958

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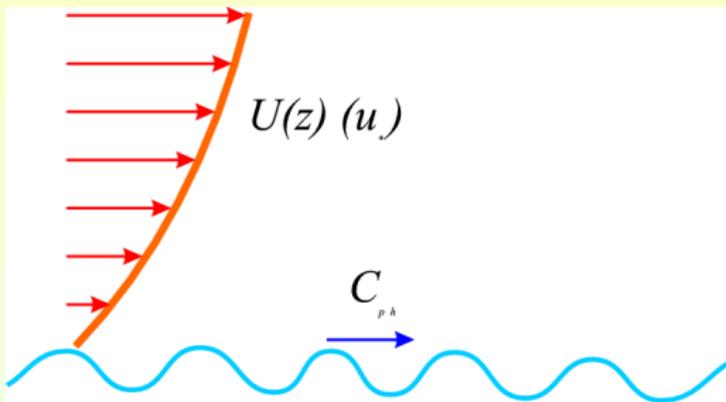
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Quasi-linea wave input

$$S_{in} = \beta(\mathbf{k}, N_{\mathbf{k}}) N_{\mathbf{k}},$$

$$\text{wave growth } \beta(\mathbf{k}) = \varrho \omega(\mathbf{k}) (\varsigma - 1)^n \quad \text{at } \varsigma > 1, n = 1, 2.$$

The Cherenkov-like factor

$$\varsigma = s \frac{U_h}{C_{ph}} \cos \theta, \quad s = O(1), \quad \theta \text{ is wave-to-wind direction.}$$



What is “true wave input”?

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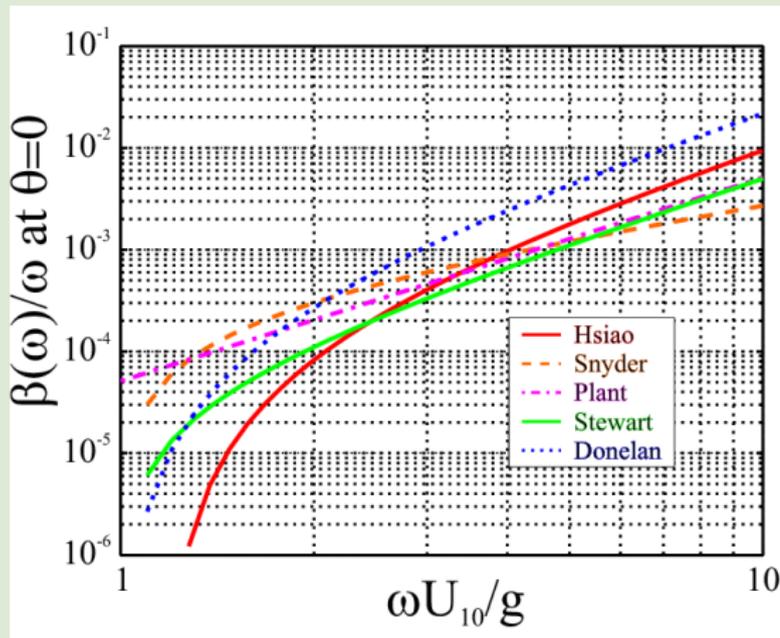
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What growth rate is correct?



Growth rate parameterizations used in wave modeling

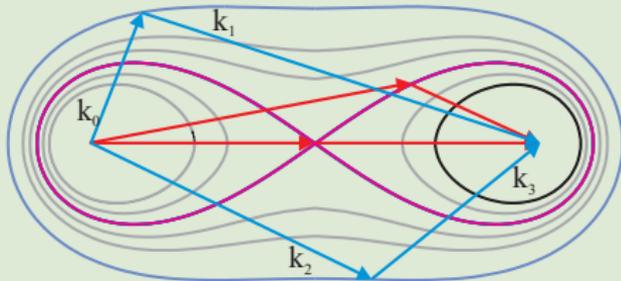


Hasselmann equation. Collision integral

$$S_{nl}[N_k] = \int_{\mathbf{k}_1, \mathbf{k}_2, \mathbf{k}_3} |T_{\mathbf{k}, \mathbf{k}_1, \mathbf{k}_2, \mathbf{k}_3}|^2 \{N_2 N_3 (N + N_1) - N N_1 (N_2 + N_3)\} \\ \times \delta(\mathbf{k} + \mathbf{k}_1 - \mathbf{k}_2 - \mathbf{k}_3) \delta(\omega + \omega_1 - \omega_2 - \omega_3) d\mathbf{k}_1 d\mathbf{k}_2 d\mathbf{k}_3$$

4-wave resonances

$$\begin{cases} \omega_0 + \omega_1 = \omega_4 + \omega_3 \\ \mathbf{k}_0 + \mathbf{k}_1 = \mathbf{k}_4 + \mathbf{k}_3 \end{cases}$$



Explicit formula for

T_{0123}

$$T_{1234} = - \frac{1}{32\pi^2} \frac{1}{(q_1 q_2 q_3 q_4)^{3/4}} \\ \times \left\{ \begin{aligned} &(-k_2 k_3 + q_2 q_3)(-k_1 k_4 + q_1 q_4) \\ &+ (-k_1 k_3 + q_1 q_3)(-k_2 k_4 + q_2 q_4) \\ &+ (k_1 k_2 + q_1 q_2)(k_3 k_4 + q_3 q_4) \\ &+ \omega_1 \omega_2 \omega_3 \omega_4 \\ &\times \left[q_1^2 + q_2^2 + q_3^2 + q_4^2 - q_{1-3}(\omega_2 - \omega_4)^2 \right. \\ &\quad \left. - q_{2-3}(\omega_2 - \omega_3)^2 - q_{1+2}(\omega_3 + \omega_4)^2 \right] \end{aligned} \right\} \\ + \frac{(\omega_2 - \omega_4)^2}{q_{1-3} - (\omega_2 - \omega_4)^2} \\ \times [2k_1 k_3 + \omega_1 \omega_3 (q_1 + q_3 - q_{1-3})] \\ \times [2k_2 k_4 + \omega_2 \omega_4 (q_2 + q_4 - q_{1-3})] \\ + \frac{(\omega_2 - \omega_3)^2}{q_{2-3} - (\omega_2 - \omega_3)^2} \\ \times [2k_1 k_4 + \omega_1 \omega_4 (q_1 + q_4 - q_{2-3})] \\ \times [2k_2 k_3 + \omega_2 \omega_3 (q_2 + q_3 - q_{2-3})] \\ + \frac{(\omega_3 + \omega_4)^2}{q_{1+2} - (\omega_3 + \omega_4)^2} \\ \times [2k_1 k_2 + \omega_1 \omega_2 (q_1 + q_2 - q_{1+2})]$$

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Nonlinear transfer in spectral wave forecasting models

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L. Cavaleri et al. / Progr. in Oceanogr. 75 (2007) 603-674

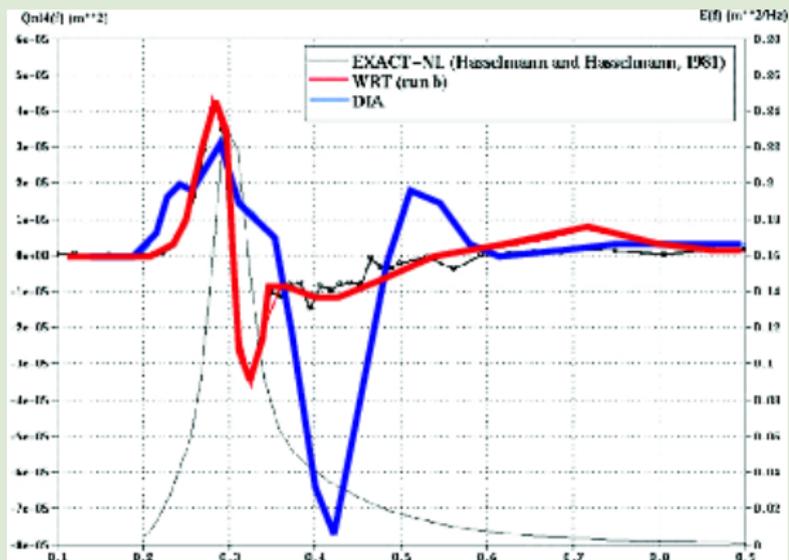
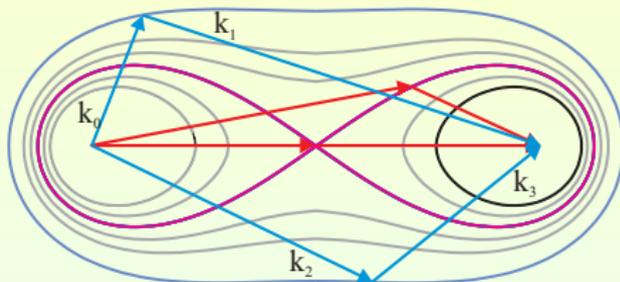


Fig. 7. Nonlinear transfer term $Qn14(f)$ computed for the test wave spectrum in deep water. These curves are obtained by integration over wave directions.



Direct Interaction Approximation for the collision integral S_{nl}

DIA takes into account just 4 (four!) symmetric quadruplets from all the continuum (!?)



The modeling within the approach requires solution of great number of equations (say, $24(\text{angles}) \times 25(\text{frequencies})$) for spatial spectra while only two characteristics of the wave field (H_s , T_s – **significant height and period**) are really used in wave forecasting.

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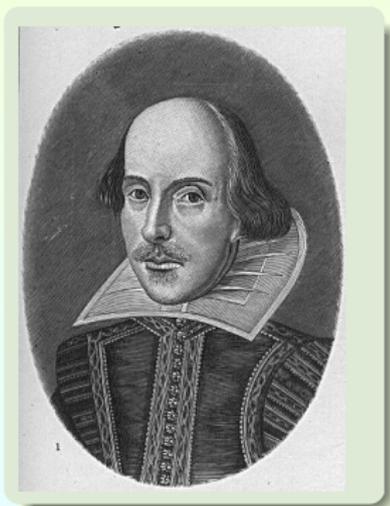
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Summary of the section

“Something is rotten
in the state of Denmark”

William Shakespeare



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A misguiding star of wind-wave modeling

by Komen, Hasselmann, Hasselmann 1984



Komen, Hasselmann, Hasselmann 1984,

“On the existence of a fully developed wind-sea spectrum”

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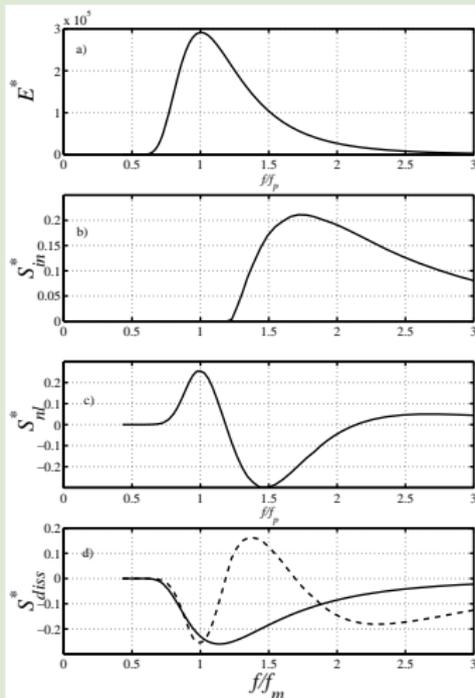
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Summary



- Pierson-Moskowitz spectrum
- Input by Snyder et al. (1981)

$$S_{in}(\omega) = \max(0, 0.25 \rho_a / \rho_w \omega \times (28 u_* / C_p - 1));$$

- Dissipation

$$S_{diss} = -S_{in} - S_{nl}$$

Output

$$\tilde{S}_{in} : \tilde{S}_{nl} : \tilde{S}_{diss} = 3 : (-1) : (-2)$$

$$\tilde{S}_i = \int_0^{2.5 f_m} S_i df d\theta$$



Q. What is responsible for wind-wave balance?

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Answers

- 1 Mainstream
Wave input and dissipation provide a
relaxation to an inherent state



Q. What is responsible for wind-wave balance?

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Answers

1 Mainstream

Wave input and dissipation provide a relaxation to an inherent state

2 Non-conventional ?

Conservative nonlinear transfer term contains both forcing and damping and is able to provide the strong relaxation on its own !!!



Nonlinear forcing and damping

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$$S_{nl} = \pi g^2 \int |T_{0123}|^2 (N_1 N_2 N_3 + N N_{\mathbf{k}_2} N_{\mathbf{k}_3} - N N_{\mathbf{k}_1} N_{\mathbf{k}_2} - N N_{\mathbf{k}_1} N_{\mathbf{k}_3}) \times \delta(\mathbf{k} + \mathbf{k}_1 - \mathbf{k}_2 - \mathbf{k}_3) \delta(\omega_{\mathbf{k}} + \omega_1 - \omega_2 - \omega_3) d\mathbf{k}_1 d\mathbf{k}_2 d\mathbf{k}_3 \quad (1)$$

Split into two terms

$$S_{nl} = F_{\mathbf{k}} - \Gamma_{\mathbf{k}} N_{\mathbf{k}} \quad (2)$$

where

$$F_{\mathbf{k}} = \pi g^2 \int |T_{0123}|^2 N_1 N_2 N_3 \times \delta(\mathbf{k} + \mathbf{k}_1 - \mathbf{k}_2 - \mathbf{k}_3) \delta(\omega_{\mathbf{k}} + \omega_1 - \omega_2 - \omega_3) d\mathbf{k}_1 d\mathbf{k}_2 d\mathbf{k}_3 \quad (3)$$

$$\Gamma_{\mathbf{k}} = \pi g^2 \int |T_{0123}|^2 (N_1 N_2 + N_1 N_3 - N_2 N_3) \times \delta(\mathbf{k} + \mathbf{k}_1 - \mathbf{k}_2 - \mathbf{k}_3) \delta(\omega_{\mathbf{k}} + \omega_1 - \omega_2 - \omega_3) d\mathbf{k}_1 d\mathbf{k}_2 d\mathbf{k}_3 \quad (4)$$



Split S_{nl} into two terms (N.N. Ivenskikh approach based on WRT-algorithm)

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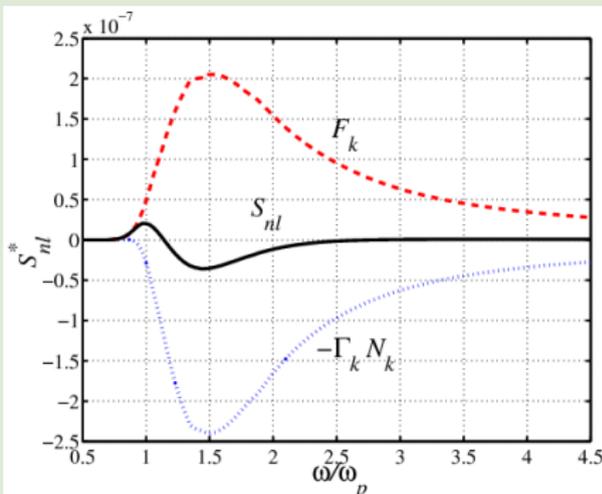
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Summary



- S_{nl} consists of two great terms of opposite signatures – forcing and damping (Hasselmann mentioned this feature);
- S_{nl} is small due to proximity to an inherent state !



Scheme N 1 of spectral balance

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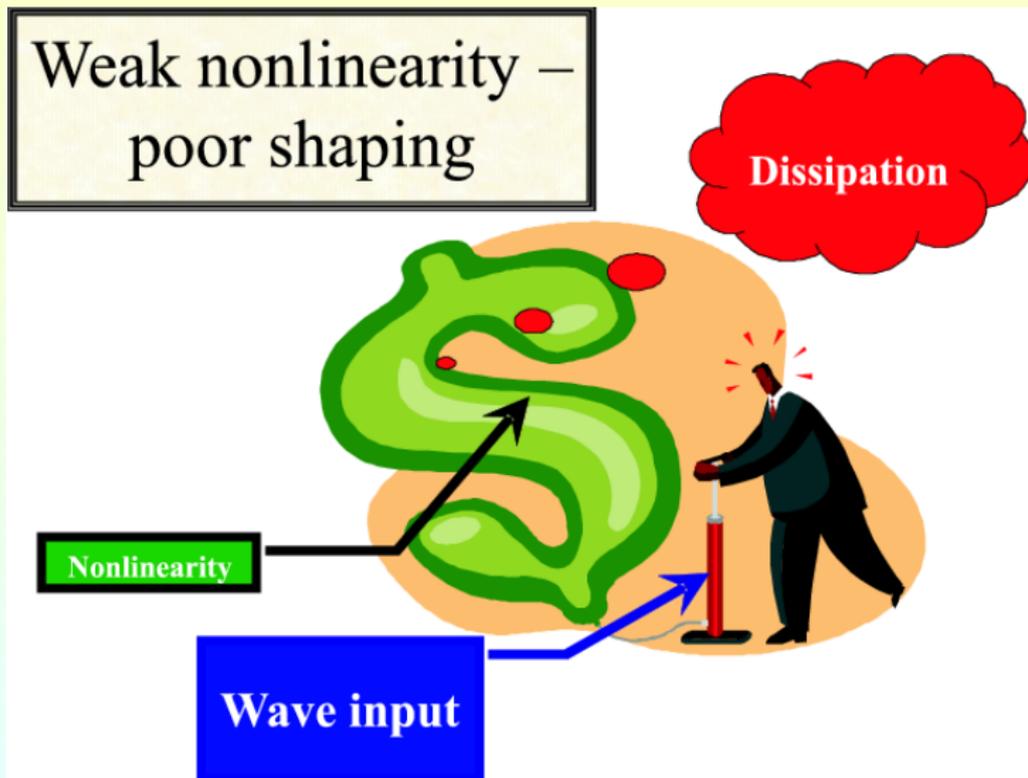
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Scheme N 2 of spectral balance

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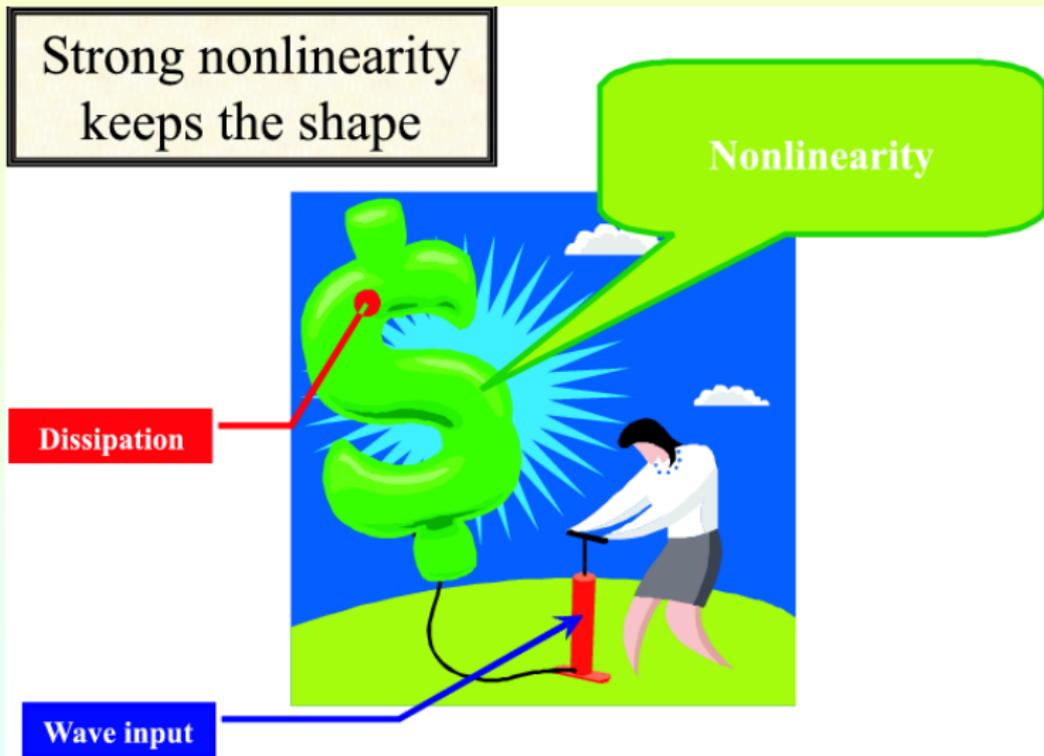
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An analytical estimate of Γ_{nl}

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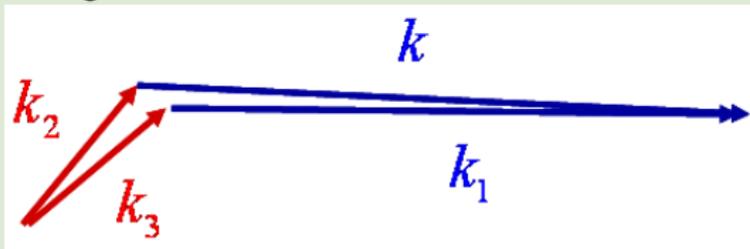
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Hypothesis: The key contribution to Γ_{nl} is from interactions of pairs of long and short waves



$$\Gamma_{\mathbf{k}} = 36\pi\omega(\omega/\omega_p)^3\mu_p^4\cos^2\Theta,$$

small parameter $\mu_p = \sqrt{\frac{E\omega_p^4}{g^2}}$ – wave steepness

An enhancing factor: $36\pi \approx 113.1$



Compare nonlinear damping decrement and wind input increment

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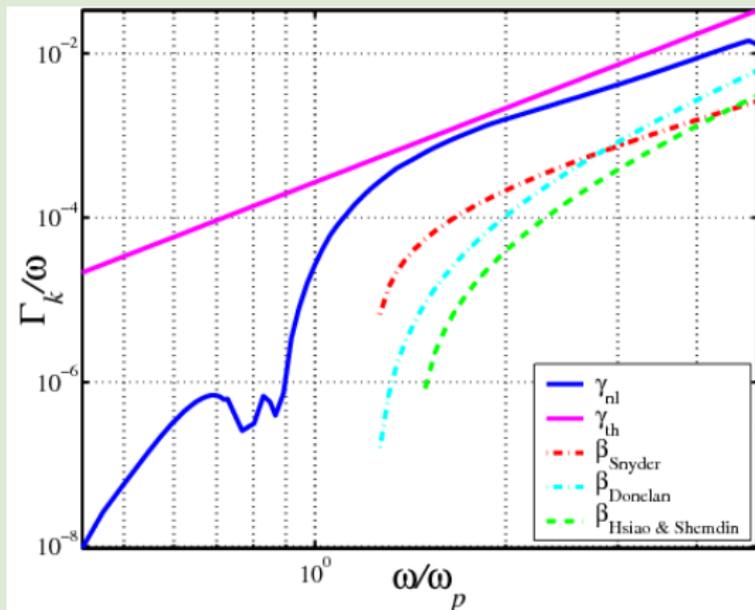
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S_{nl} surpasses S_{in} and S_{diss} in order of magnitude !



Mixed sea

I. R. Young, 2006, JGR

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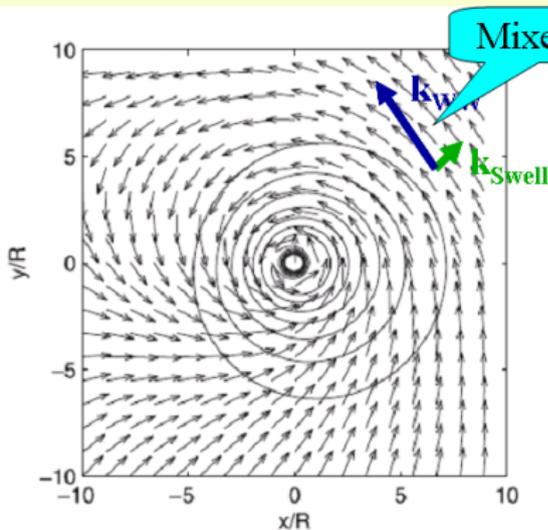


Figure 4b. A typical wind field for a Northern Hemisphere hurricane, as predicted by the model (1). The contours are of wind speed, U_{10} , with the vectors showing wind direction. The scale is identical to Figure 4a. The system is shown for the Northern Hemisphere (i.e. anti-clockwise circulation).

YOUNG: HURRICANE DIRECTIONAL SPECTRA

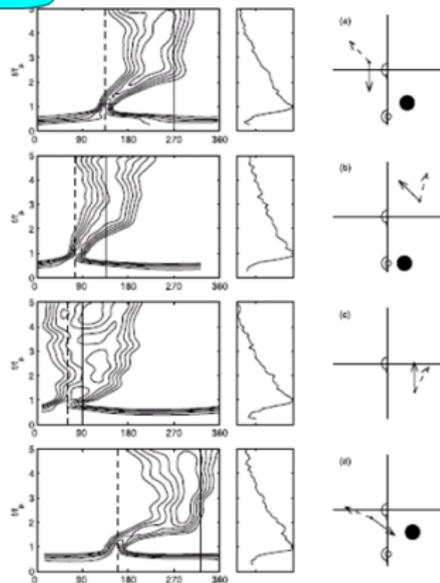


Figure 5. Examples of the directional spreading function, $D(f, \theta)$ for each quadrant of a hurricane.



Outcome:

I. R. Young, 2006, JGR

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The non-linear wave-wave interaction term
is actually stronger than the representations
which are implemented in,
even the most sophisticated, research models



“The least sophisticated” model of the mixed sea (WISE-2008)

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Summary

Nothing special: solve numerically spatially homogeneous kinetic equation (the Hasselmann equation)

$$\frac{\partial N}{\partial t} = S_{nl} + S_{in} + S_{diss}$$

with exact nonlinear transfer term S_{nl}



Case 2 by Young. Close directions of swell and wind waves. Simulations

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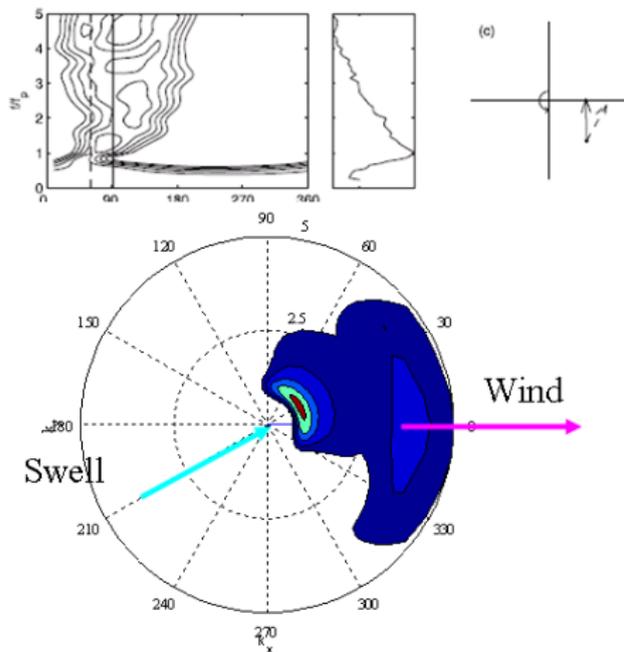
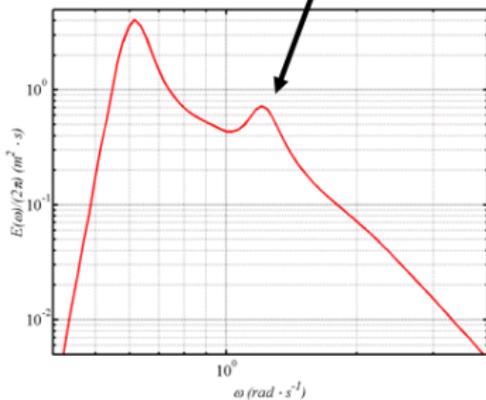
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$$U_{10}=20\text{m/sec}; H_s=8\text{m};$$
$$f_p=0.09\text{ Hz}; f_{2p}=0.192\text{Hz}$$





Case 2 by Young. Close directions of swell and wind waves. Simulations

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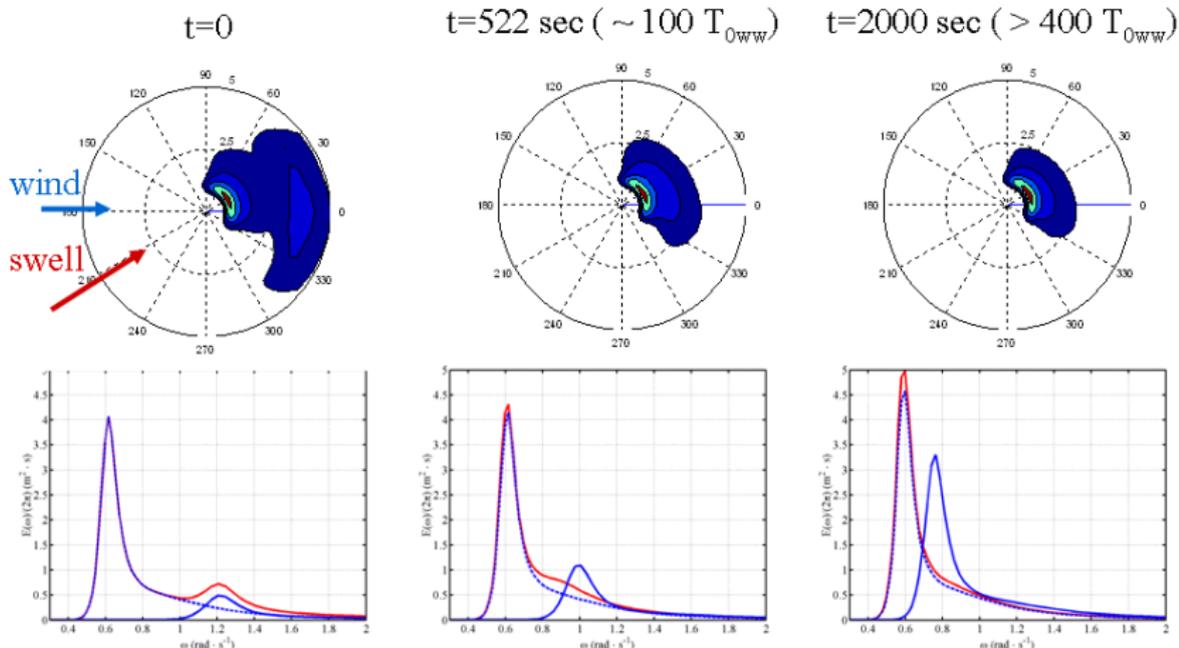
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Steering, absorbing, reduction of direct wind input *etc.*



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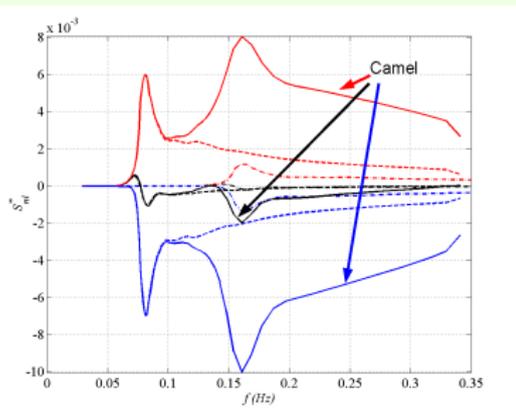
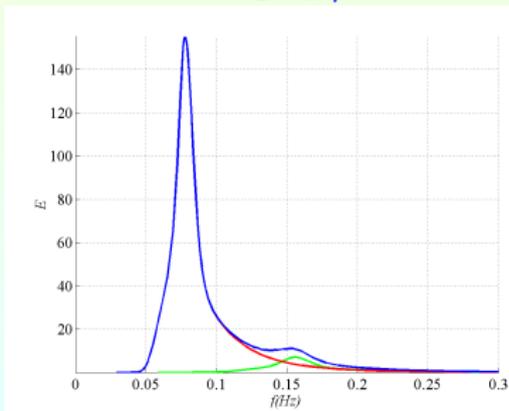
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Summary

Can we explain the fast transition within our simple estimate of nonlinear damping?

$$\Gamma_k = 36\pi\omega(\omega/\omega_p)^3\mu_p^4\cos^2\Theta,$$

Explanation: High (ω/ω_p) – from swell peak,
high μ_p – from wind waves





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Summary

- 1 **Nonlinear relaxation**, generally, is much stronger than quasi-linear external forcing and wave dissipation;



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Summary

- 1 **Nonlinear relaxation**, generally, is much stronger than quasi-linear external forcing and wave dissipation;
- 2 **Interactions of long and short waves** play key role in this relaxation;



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Summary

- 1 **Nonlinear relaxation**, generally, is much stronger than quasi-linear external forcing and wave dissipation;
- 2 **Interactions of long and short waves** play key role in this relaxation;
- 3 **We do not ignore wave input and dissipation**, we just put them into proper place